

Research Article

The characteristics and genesis of volcanic ash soils at the northern slope toposequence of Kawi Mountain in Malang Regency

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Abstract: The volcanic ash soil in Kawi Mountain is composed by the amorphous materials consist of allophane, imogolite and ferrihydrite. Results of previous study showed that the phosphate retention in all soil profiles of northern slope toposequence of Kawimountain was less than 85%, yet the phosphate retention of volcanic ash soils is usually $\geq 85\%$. This raised a question that there is a different characteristics of soil in the northern slope tosequence of the Kawi Mountain compared to the other places. This research was conducted to study soil characteristics, mineral contents, and genesis processessoccurring in soil on the northern slope toposequence of Kawi Mountain. 5 pedons between high elevation and low elevation (P1, P2, P3, P4 and P5) to identified the soil characteristics. The Al, Fe, and Si extracted by acid oxalate, natrium pyrophosphate, and dithionite citrate to calculate the amorphous mineral content. The results show that. The results showed that there is a different in terms of the thickness of the A horizon, the C organic content and the soil acidity level that mainly found in P3 and P4 profiles. The most important soil genesis processess in the formation of the volcanic ash soils were likely clay illuviation (P5), melanization and braunification (P3), littering (P1) and the reduction of andic soil properties from the upper slope (P1 profile) up to the lower slope (P5 profile).

Keywords: *genesis, Kawi Mountain, toposequence, volcanic ash soil*

Introduction

Volcanic ash soil developed from volcanic activity materials such as tephra or pyroclastic (Shoji et al., 1993). According to Leamy et al. (1988), the earth's land surface covered by volcanic ash soils about 124 milion hectare. Soils developed from the volcanic materials usually classified as Andisols that have an unique characteristics such as low bulk density ($\leq 0.9 \text{ g.cm}^{-3}$), high $\text{Al}_0+1/2 \text{ Fe}_0 \geq 2\%$ and high phosphate retention ($>85\%$). Those characteristics are called andic soil properties.

Based on the Team Nuffic-Unibraw (1984), Kawi Mountain (2631 masl) is a stratovolcano that extend outward in small shear from the north to south and consists of a combination of andesitic rock layers, breccias, agglomerate, tuff (lava material and harden pyroclastic) and ash that overlapping on each other. Kawimountain has a

various topography that effect to the soil characteristics and genesis in the northern slope of Kawi Mountain. In per-survey stage, we found that the andic soil properties especially in phosphate retention and Si (extracted be acid oxalate) are a very low. This condition brings up a question "why the volcanic ash soil different with the same soil in another places". This study was conducted to explain the characteristics and genesis volcanic ash soil on northern slope Kawi Mountain. This paper also examine the mineralogy contents in northern slope toposequence of Kawi Mountain in Malang Regency.

Material and Method

This study was done in the northern slope toposequence of Kawi Mountain from started

June 2013 to August 2014. The Kawi Mountain administratively lies in the Bendosari Village, Pujon District, Malang Regency, East Java. The coordinate of the study location is shown in Figure 1. Observation of soil morphological characteristics and collection of soil sample from each soil horizon were conducted in the profile of 180 cm depth. The collected soil samples were analyzed in a laboratory to determine the texture, the cation exchange capacity, the base saturation,

Organic, water retention at 1500 kPa, pH, as well as Al, Si, Fe (oxalate, pyrophosphate and dithionite) contents. Soil bulk density was determined for undisturbed soil samples heated up to 105°C. The 1500 kPa water retention was measured using pressure plate through the gravimetric method. The soil pH was measured using the pH electrode meter in which the prepared sample was added with water (pH H₂O) or KCl (pH KCl) with 1:1 proportion.

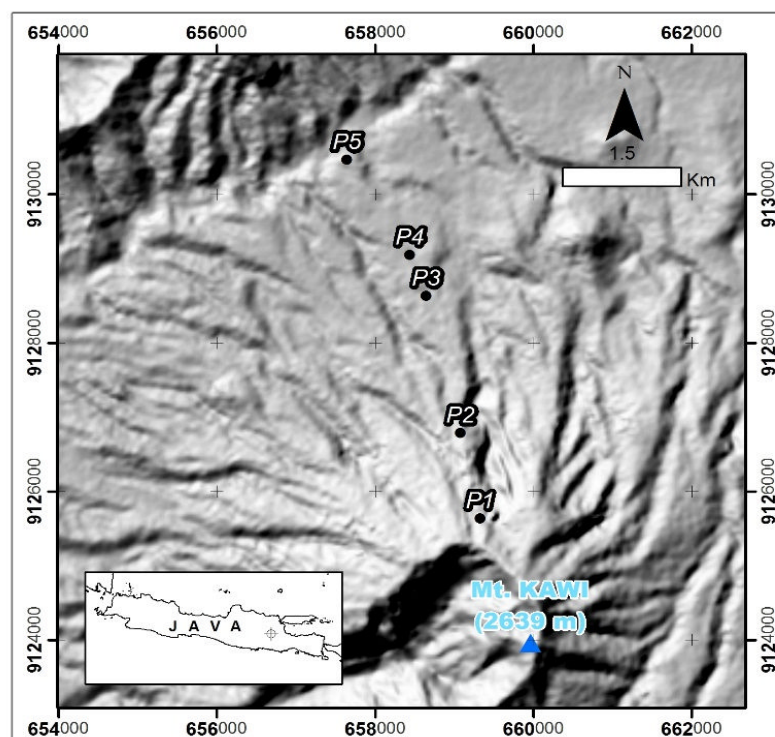


Figure 1. The toposequence scheme of the northern slope of Kawi Mountain. P1 = Pedon 1, 35°, 2150 m (Typic Hapludands), P2 = Pedon 2, 35°, 1610 m (Typic Hapludands), P3 = Pedon 3, 26°, 1195 m (Humic Udivitrands), P4 = Pedon 4, 14°, 1149 m (Andic Dystrudepts), P5 = Pedon 5, 5°, 1061 m (Ruptic-Alfic Dystrudepts)

The C-organic was determined using the Walkley & Black method. The cation exchange capacity (CEC) and exchangeable Ca, Mg, Na and K were extracted by using acetate ammonium pH 7. The base saturation (BS) was counted by using an equation of $BS (\%) = (\sum Ca, Mg, Na, K) / CEC \times 100 \%$ (ISRIC, 1993).

The active Al, Si, and Fe were extracted using ammonium oxalate while the bounded Al, Si and Fe were extracted with the sodium pyrophosphate (Blackmore et al., 1987), and the free Al, Si, and Fe were extracted with the dithionite citrate (Mehra and Jackson, 1960). Phosphate retention was measured by the method of Blackmore et al. (1987). Allophane and imogilite were counted using an equation of $Si_o \times 7.14$ (Shoji et al., 1988). Ferrihydrite was

measured using an equation of $Fe_o \times 1.7$ (Mizota and van Reeuwijk, 1989). The crystallization degree was counted using Fe_o/Fe_d equation (Schwertmann, 1985). The formed of allophane and Al-humus were measured with Al_p/Al_o equation (Parfitt, 1985).

Results and Discussion

Soil morphological characteristics

The soils in the study area generally have A/Ap-AB-Bw-BC-2BA/2Bw horizons. The A/Ap horizon thickness ranges from 20 cm (P4) to 43 cm (P3). The AB horizon transition is an A horizon which has a little of B horizon characteristic and under this horizon there are Bw

and BC horizons. In P3, P4 and P5 profiles, under the Bw/BC there is lithologic discontinuity which symbolized with 2BA/2Bw. The clearness of the horizon usually looks obvious and the boundary is flat though, in some horizons, they look clear and gradual with the wavy boundary. Soils in the P1 and P3 profiles tend to have dark colour in the upper surface and it gets brighter in the lower one. The contrast of the bright colour has chrome from 4 to 6 (Table 1).

Soils in the P2 and P5 profile have a less changing colour in both upper and lower horizons. The whole profiles tend to be dark whenever the amount of organic material is high. The intensive use of soil in the P2 location can cause the decrease of the organic material content which will affect the soil colour. The soil texture in the P1, P2 and P3 profiles is dominated by sand and silt fractions. The upper surface has sandy clay texture and it changes to be silty clay in the lower surface. The texture in upper profiles of P4 and P5 is clay and it changes to be firm clay in the lower surface. Soils in the study area have a low bulk density value, except that of the P5 profile.

When it is related to the requirements of andic soil properties (the bulk density is $< 0.9 \text{ g/cm}^3$), soils of the P1 to P4 profile have already fulfilled the requirement that is the lower the horizon, the lower is the bulk density value. The soil bulk density in the P5 profile ranges from 1.2 to 1.3 g/cm^3 . The bulk density values observed in this study is in the opposite of the percentage of C-organic, the deeper is the profile, the higher is the content of C-organic and the bulk density tends to be lower. The water retention ranges from 10 to 30% in all profiles and it tends to be higher in the lower horizon compared to the upper one. The physical characteristics of the soil is presented in Table 1.

Soil chemical characteristics

The pH (H_2O) of soils in the study area is less than 7 which tends to be higher in the deeper soil profile. The lowest soil pH (4.76) is in the upper horizon of P5 profile (Table 2). According to Shoji and Fujiwara (1984), when the soil pH is 5 or less then the organic acid is dominated with proton donor in order to form Al-humus since it is in volcanic ash soil. Therefore, the allophane or other amorphous materials will be formed in pH 5. The frequent tendency is that the higher value of the pH, the higher is the amount of the amorphous minerals.

Although the value of pH KCl is relatively stable in the whole horizons, the value of the pH KCl is smaller than the pH H_2O , except that in the P5 profile. The pH of NaF can only be measured

when it is in 36 cm depth from the soil surface. P1, P2 and P3 profiles have a high value (more than 9.2) and the P4 profile has less than 9.2. According to Msanya (2007), the P retention indicates the amount of the hydroxyl cluster-Al(Fe) that have reacted with NaF which in turn causes the loss of OH^- from the bound. The P retention is mostly affected by the existence of amorphous minerals such as allophane and imogolite in the soil.

The cation exchange capacity (CEC) in P1 ranges from 47.13 cmol/kg in the upper horizon to 73.15 cmol/kg in the lower surface horizon. In the P2, P3 and P5 profiles, the CEC value tends to be stable that ranges from $>18 \text{ cmol/kg}$ to 35 cmol/kg. In the P4 profile, the CEC value ranges from 46 to 61 cmol/kg. The Base saturation (BS) in P1 to P4 profiles is less than 35%. The BS value in the P5 profile which is $>35\%$ in all horizons, is higher than that of other profiles. The C-organic content in the study area is high in P1 up to P3 profiles (Table 2). The C-organic content in the P1 and P2 profiles ranges from 2.88% to 5%. In the P4 and P5 profiles, however, value of C-organic is low. The average value of C-organic in P4 and P5 profiles is less than 1%.

The different elevation height of pedon location shows that the higher the place, the higher is the value of the phosphate retention. In P5 to P3 profiles, the high value of phosphate retention is no longer found. The range of value is from 10 to 60 %. However, the phosphate retention in P2 and P1 profiles can reach up to 80%. In the area of soil with volcanic ash parent material, the affinity of Al and Fe toward phosphate is high. So, it is assumed that the closer to the parent material source, the higher is the value of phosphate retention. $\text{Al}+1/2 \text{ Fe}$ extracted with the oxalate acid is one of the criteria of andic soil. If the value of $\text{Al}+1/2 \text{ Fe}$ is 2 or more, it is strongly assumed that the place has the andic soil properties. Almost all profiles and horizons have the value of $\text{Al}+1/2 \text{ Fe}$ more than 2, except for the P1 horizon 1.

Soil Mineral Content

The extraction of Ammonium Oxalate and Al is around 2.03% in the upper slope and 0.2 – 6% in the lower slope. Based on Parfitt *et al.* (1984) and Shoji *et al.* (1993) the concentration of the Al_0 in Andisol is various, from 0.02 – 4.13%. Extraction that done toward Fe in the upper slope is around 0.87 up to 1.57%. Based on Parfitt *et al.* (1984) and Shoji *et al.* (1993), the value of Fe is at the range of 0.02 – 0.87%. The result of the ammonium oxalate extraction is shown in Table 2.

Table 1. Physical characteristics of soils in the northern slope of Kawi Mountain

Pedon	Horizon	Depth (cm)	Colour		Texture ¹	Structure ²	Consistency ³	Horizon Boundary	Bulk Density (g/cm ³)	Water retention (1500 kPa)
			Wet	Dry						
Pedon 1	A	0-34	10 YR 3/2	10 YR 3/2	SL	m-m-ab/g	f/s-ra-np	Clear- Smooth	0.75	15.52
	AB	34-71	10 YR 4/6	10 YR 3/4	SL	m-m-ab/sab	f/rs-ra-np	Clear- Smooth	0.65	25.55
	BA	71-98	10 YR 3/6	10 YR 4/3	SiL	m-m-ab/sab	f/rs-ra-np	Clear- Smooth	0.60	21.59
	Bw1	98-113	10 YR 3/6	10 YR 4/4	SiL	m-m-ab/sab	f/rs-ra-np	Clear- Smooth	0.43	25.21
	Bw2	113-141	10 YR 4/6	10 YR 5/4	SiL	m-m-ab/sab	f/rs-ra-np	Clear- Smooth	0.50	34.87
Pedon 2	A	0-29	10YR 2/2	10 YR 3/3	SL	m-m-g	s-ra-rp	Clear- Smooth	0.64	20.71
	AB	29-73	10 YR 2/1	10 YR 2/2	SL	m-m-sab/ab	s-ra-rp	Clear- Smooth	0.71	25.22
	Bw1	73-95	10 YR 2/1	10 YR 3/2	SiL	m-m- sab/ab	f/rs-ra-np	Clear- Smooth	0.62	28.65
	Bw2	95-150	10 YR 2/1	10 YR 3/2	SiL	m-m- sab/ab	f/rs-ra-np	Clear- Smooth	0.52	25.31
Pedon 3	A	0-43	10 YR 2/2	10 YR 4/2	SL	m-w-sab	f-ra-rp	Clear- Smooth	0.88	12.10
	AB	43-66	10 YR 2/1	10 YR 3/2	SiL	m-m-sab	rs-ra-rp	Clear- Smooth	0.68	17.92
	2BA	66-100	10 YR 3/3	10 YR 6/4	SL	m-m-sab	s-ra-rp	Clear- Smooth	0.89	15.97
	2Bw1	100-124	10 YR ¾	10 YR 6/3	SiL	m-m-ab	s-ra-rp	Clear- Smooth	0.99	16.52
	2Bw2	124-141	10 YR ¾	10 YR 5/4	SiL	m-m-sab	rs-ra-rp	Clear-Wavy	0.96	(-)
	2Bw3	141-170	10 YR 3/6	10 YR 5/4	SiL	m-m-sab	f-ra-rp	Clear-Wavy	0.84	(-)
Pedon 4	Ap	0-20	10 YR 4/6	10 YR 5/4	L	m-m-g	f-ra-np	Gradual- Smooth	0.85	24.38
	Bw1	20-39	10 YR ¾	10 YR 5/4	CIL	m-m-g	rs-a-p	Clear- Smooth	0.70	27.92
	Bw2	39-76	7,5 YR 2,5/2	10 YR 2/2	CIL	m-m-sab	rs-a-p	Clear- Smooth	0.83	33.18
	2Bw3	76-103	7,5 YR 3/3	7,5 YR 3/2	L	f-m-sab	rs-ra-rp	Clear- Smooth	0.53	(-)
	2Bw4	103-140	7,5 YR 2,5/1	7,5 YR 3/2	L	f-m-sab	rs-ra-rp	Clear- Smooth	0.47	(-)
Pedon 5	A	0-31	7,5 YR 3/2	10 YR 6/2	L	m-m-sab/g	f-ra-rp	Gradual- Smooth	1.25	25.43
	AB	31-59	7,5 YR 3/2	10 YR 6/2	L	m-m-sab/g	rs-ra-rp	Clear-Wavy	1.36	26.58
	Bw1	59-77	7,5 YR 3/3 - 10 YR 3/3	10 YR 6/2	SL	m-m-ab/sab	rs-ra-rp	Clear- Smooth	1.30	(-)
	BC	77-85	7,5 YR 3/3 - 10 YR 5/4	10 YR 6/2	SL	m-m-ab/g	s-ra-np	Abrupt- Smooth	1.30	(-)
	2Bw2	85-117	7,5 YR 3/2	10 YR 5/4	CIL	m-m-ab/sab	rs-a-p	Clear- Smooth	1.23	(-)
	2Bw3	117-132	10 YR 3/3	10 YR 5/4	CIL	m-m-ab/sab	rs-ra-rp	Clear- Smooth	1.20	(-)

¹LS: Loamy Sand, SiL: Silty Loam, L: Loam, CIL: Clay Loam

²Size (f: fine, m: medium, c: coarse),Development (w: weak, m: medium, h: hard), Type (g: granular, sab:sub angular blocky, ab: angular blocky)

³f: friable, rs: rather shady, s: shady, ra: rather adhesive,a: adhesive, rp: rather plastic, p: plastic, np: no plastic

Table 2. Chemical characteristics of soils in the northern slope of Kawi Mountain

Pedon	Horizon	pH		Cation Exchange Capacity (cmol/kg)	Base Saturation (%)	C-organic (%)
		H ₂ O	KCl			
Pedon 1	A	5.93	5.28	47.13	23.02	3.54
	AB	6.24	5.68	57.47	27.68	3.28
	BA	6.17	5.67	59.09	11.43	2.95
	Bw1	6.24	5.49	73.15	9.03	3.78
	Bw2	6.25	5.56	57.18	13.64	2.88
Pedon 2	A	6.01	5.3	36.50	24.15	3.84
	AB	6.20	5.47	38.30	19.27	4.68
	Bw1	6.13	5.43	47.23	15.21	5.03
	Bw2	6.11	5.46	59.57	16.27	5.41
Pedon 3	A	5.98	4.87	34.82	23.05	3.86
	AB	5.91	5.04	32.53	28.65	4.10
	2BA	6.18	5.44	19.24	34.76	1.65
	2Bw1	6.07	5.07	35.88	19.98	1.38
	2Bw2	6.16	5.14	31.62	25.52	1.55
	2Bw3	6.14	5.28	18.58	33.47	0.55
Pedon 4	Ap	5.38	4.08	56.38	18.42	1.34
	Bw1	5.84	4.79	46.95	14.69	0.79
	Bw2	5.94	5.15	62.03	19.68	2.59
	2Bw3	5.87	4.69	61.77	17.53	1.11
	2Bw4	6.09	5.18	58.80	20.61	1.58
Pedon 5	A	4.76	5.32	30.81	35.17	0.94
	AB	4.92	5.55	25.23	41.23	0.54
	Bw1	4.98	5.85	24.60	41.18	0.21
	BC	4.98	5.85	24.60	41.18	0.21
	2Bw2	5.00	5.97	24.89	41.03	0.14
	2Bw3	5.03	6.19	30.64	40.30	0.29

This condition describes that the value of Fe in the further research is higher compare to the range given by Parfit and Shoji. The Si value that was extracted with the ammonium oxalate is around 0.03 up to 0.1 % in the upper slope and 0.1% to 0.2 % in the lower slope. Those numbers are considered as a small number if they are compared to what has found by Tsai (2010) that is 0.82 – 1.4% in the upper slope and 0.1 – 0.6% in the lower slope. The extraction result of Pyrophosphate to Al in the upper slope is 0.02% while in the lower slope is 0.04%. The value of Pyrophosphate extraction with Fe in the upper slope is 0.3 up to 0.4% and 0.4 to 0.8% in the lower one. The Dithionite extraction value with Al is around 0.86 to 1.35 % in the upper slope and 0.2 to 0.6 % in the lower slope.

Based on Schwertmann (1985), the ratio of Fe_o and Fe_d can be used as the indicator of the crystallization degree of the iron oxide age and it also can be reflected as the soil development degree of volcanic ash soils (Malucelli *et al.*, 1999). Kleber *et al.* (2004) says that the volcanic ash soil (mainly Andisol) has the Fe_o and Fe_d ratio more than 0.75. In P1 Fe_o/Fe_d more than 0.75 between 0 until 98 cm depth and decreasing

in the lower horizon. Pedon P2 and P3 have no different condition with P1 Fe_o and Fe_d ratio. The Fe_o and Fe_d ratio shows that crystallization degree in P4 and P5 is higher than P1, P2 and P3 soil profiles. It means that the soil in P4 and P5 have more develops than the soil in the upper slope (P1, P2 and P3). The result of Fe_o/Fe_d shown in Table 4.

All Pedon in each horizon has Al_o/Al_p value less than 0.5. This condition can be reflected that no Al-humus bound was found in the soil. The reason of inexistence the Al-humus complexes caused the pH H₂O more than 5 and it support to the production of amorphous mineral in the soil. Amorphous minerals formed in the pH range of 5-7 and in the udic soil moisture regime where the aluminum hydroxyl cation reacts to form the amorphous minerals. There is no tendency that shows the dynamic of the amorphous minerals in all pedon. But, if we see the number of the most amounts of amorphous (Allophane, Imogolite and Ferrihydrite) are in P3 and P4 profiles. Generally, the higher number of Allophane and Imogolite in the Andisol was linear with the soil depth, but it does not happen in the research location.

Table 3. The analysis result on extraction of Al, Si and Fe using ammonium oxalate, pyrophosphate and dithionite

Pedon	Horizon	Oxalate Extraction			Pyrophosphate Extraction		Dithionite Extraction			Fe _o /Fe _d	Al _o +1/2 Fe _o	Amorph
		Si _o (%)	Al _o (%)	Fe _o (%)	Al _p (%)	Fe _p (%)	Si _d (%)	Al _d (%)	Fe _d (%)			
Pedon 1	A	0,01	2,03	0,62	0,02	0,47	0,25	0,79	0,65	0,95	2,34	1,12
	AB	0,05	3,85	1,46	0,02	0,42	0,32	0,88	1,36	1,07	4,58	2,84
	BA	0,04	3,94	1,49	0,02	0,36	0,39	0,86	1,80	0,83	4,69	2,82
	Bw1	0,03	4,23	1,64	0,02	0,38	0,56	1,35	3,05	0,54	5,05	3,00
Pedon 2	A	0,01	2,14	0,96	0,02	0,48	0,27	0,81	0,70	1,37	2,62	1,70
	AB	0,01	3,25	0,87	0,02	0,42	0,28	0,79	1,05	0,83	3,69	1,55
	Bw1	0,01	3,95	1,26	0,02	0,45	0,35	0,93	1,63	0,77	4,58	2,21
	Bw2	0,01	4,25	1,57	0,02	0,47	0,42	1,03	2,07	0,76	5,04	2,74
Pedon 3	A	0,45	3,41	1,32	0,02	0,29	0,26	0,48	0,94	1,40	4,07	5,46
	AB	0,08	3,88	1,50	0,03	0,44	0,32	0,53	1,02	1,47	4,63	3,12
Pedon 4	Ap	0,10	1,16	2,64	0,04	0,40	0,48	0,44	3,01	0,88	2,48	5,20
	Bw1	0,18	0,82	1,84	0,01	0,59	0,51	0,36	3,12	0,59	1,74	4,41
	Bw2	0,25	1,56	1,97	0,01	0,83	0,72	0,62	2,81	0,70	2,55	5,13
Pedon 5	A	0,14	6,00	1,00	0,01	0,45	0,33	0,26	1,63	0,61	6,50	2,70
	AB	0,14	0,52	0,92	0,01	0,45	0,33	0,20	1,52	0,61	0,98	2,56

Soil Genesis

Andisols formed in pedon P1, P2, and P3 affected by amorphous materials. Typic Hapludands that discovered in P1 have an umbric epipedon and cambic endopedon. Overall horizons has low bulk density less than 0.9 g.cm⁻³, P retention more than 25 % and Al+1/2 Fe more than 0.4 %. Pedon P2 classified as HumicUdivitrands with water retention in 1500kPa less than 15% in the top of the horizon.

Table 4. The P retention and bulk density of the soil profiles in the research location

Pedon	Horizon	P retention (%)	Bulk Density (g/cm ³)
Pedon 1	A	68,00	0,75
	AB	83,80	0,65
	BA	83,20	0,60
	Bw1	90,20	0,43
Pedon 2	A	72,70	0,64
	AB	79,50	0,71
	Bw1	81,60	0,62
	Bw2	81,90	0,52
Pedon 3	A	51,50	0,88
	AB	50,70	0,68
Pedon 4	Ap	41,10	0,85
	Bw1	28,30	0,70
	Bw2	45,40	0,83
Pedon 5	A	22,90	1,29
	AB	14,30	1,28

P4 and P5 are classified as the Inceptisol that have a low characteristics of andic soil properties. The low count of andic soil properties linear with decreasing the elevation. The lowest P retention value was found in P5 or lower elevation (P retention requirement was more than 25 % in coarse particle and 85 % in smooth particle) and the highest bulk density was in P5 to (the requirement of andic soil properties was less than 0.9 g/cm³).

Braunification is one of the soil genesis process that connected with the forming of soil colour. Based on Chen *et al.* (1999), Braunification is a process of the brown colour formation in the soil as the secondary result of the Fe Oxide formation and it can be proved by the value of Fe_d and Fe_o. Braunification mostly happens in the B horizon since in the upper surface of the soil there is a pedogenesis process that will lead to Melanization process. In P4 profile the braunification occurs in the Bw2 and 2Bw3 horizon, while in the below horizons the braunification does not happen since in those areas are still affected by the black colouring formation from the old soil. In P4 profile, the bright colour in the soil surface mainly caused by the agricultural activity as the soil utilization. This fact can be proved by the existence of the black colour soil in the deeper level. The easiest observation toward braunification process occurs in P1 and P3 profiles where the black colour in

the upper horizon changes into light brown which makes the clear boundary.

Littering is piles of fresh organic matter that only exist in P1 soil profile (the highest pedon) which is 9 cm in thickness. While in the other profiles, there is only small amount of littering accumulation. The melanization process is signed by the dark colour changing in the soil that caused by the mix between the organic matter/ humus with the mineral deposit in horizon O and A. Besides Umbric and Mollic, the result of melanization process is the forming of Melanic. Melanic has high content of C_{organic} which also have the andic soil properties in it. By seeing the value and chrome colour level (in the moist was ≤ 2) in P2 and P3 profile, the research result shows that there is a possibility of the melanic is formed. Though, that particular melanic characteristics cannot be fulfilled since the C_{organic} value is less than 4%.

Soil is classified in to Typic Hapludands in P1 and P2 profile, Humic Udivitrands in P3 profile, Andic Dystrudepts in P4 profile and Ruptic-Alfic Dystrudepts in P5 profile. Humic Udivitrands is characterized by the value of the water retention is 12.10% in A horizon ($\leq 15\%$). Inceptisol in P5 profile is stated as the most developed inceptisol since the andic soil properties in this profile is the weakest one, which showed by the value of the bulk density is more than 1 g.cm^{-3} and the phosphate retention was very low. Ruptic-Alfic Dystrudepts is described as soil that in one of its horizon has base saturation less than 35% and there is a small amount of solid illuviation in its B horizon. This condition also indicates that the farther from the parent material source, the soil will more highly develop.

Conclusion

Soil developed under volcanic ash materials on the northern slope toposequence of Kawi Mount are classified as Andisols (in the upper part) and Inceptisols (in the lower part). The different characteristics that is observable in each profile is the thickness of the A horizon, the soil C_{organic} and the soil acid. The allophane and imogolite condition that commonly getting higher whenever the deeper the soil depth is found to have an opposite result, while the ferrihydrite is getting higher when the depth of the soil is deeper. The Fe degree of crystallization in the research location is getting higher in the lower depth of the location. In P5 profile which has already contained solid illuviation while P3 profile have the melanization process that happens in 43 cm depth and also braunification under this horizon. P1

profile there is littering of organic matter 9 cm in thickness and the andic soil characteristics does exist.

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